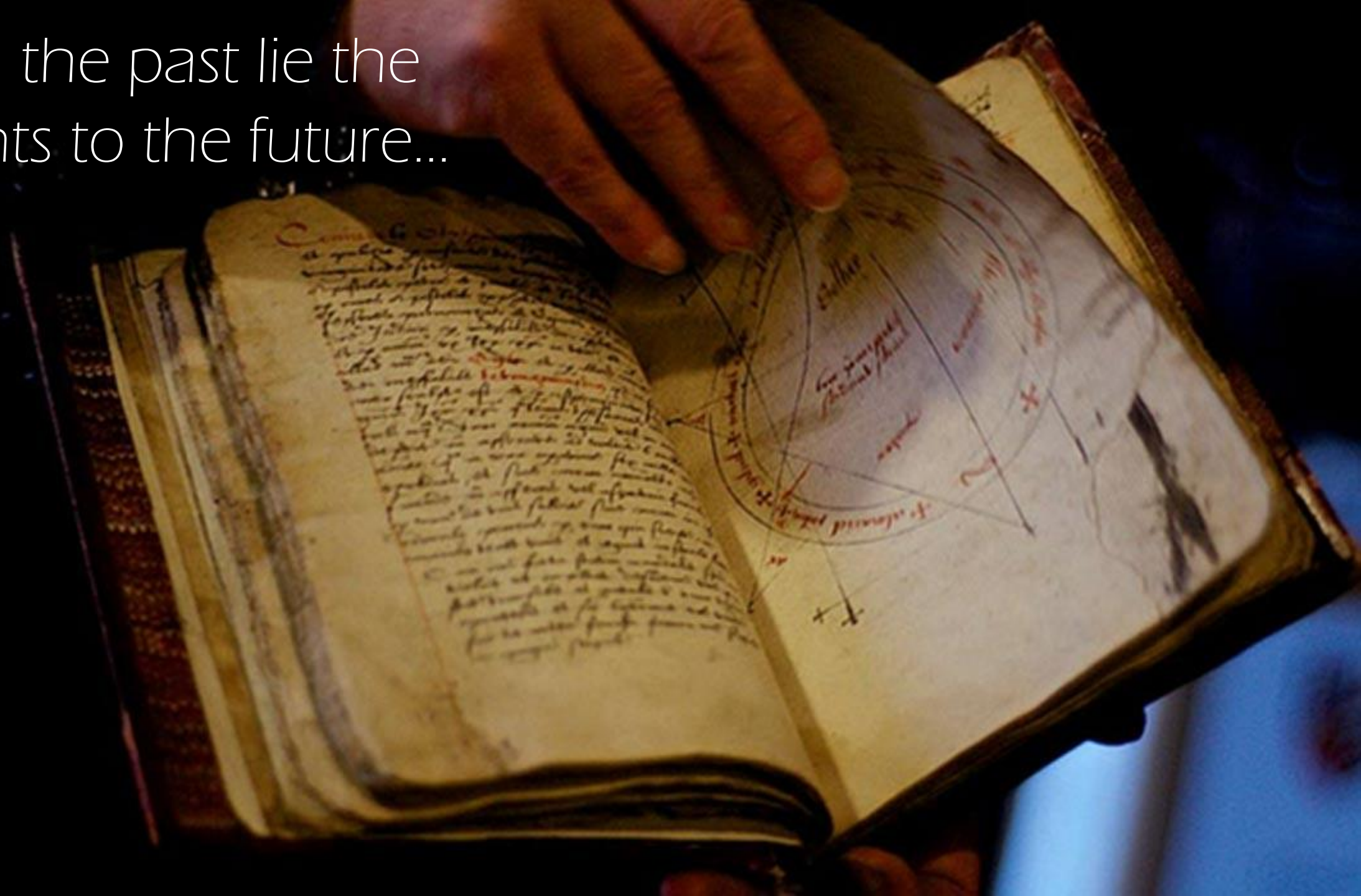


Geoneutrinos: lessons from the past, insights for the future



...within the past lie the
footprints to the future...



...1953...



In 1953 G. Gamow wrote to F. Reines: "It just occurred to me that your background may just be coming from high energy beta-decaying members of U and Th families in the crust of the Earth."

Dear Fred,
Just occurred to me that your background neutrinos my just be coming from high energy β -decaying members of U and Th families in the crust of the Earth. I do not have on the train any inform. to check it up, but it seems the order of magn. is reasonable. In fact the total energy radioactive energy production under one square foot of surface may well be equal to the energy of solar radiation falling on ~~Earth~~ that surface. ;
What do you think ;
write to me at : The Union Univ. of Mich. Ann Arbor. Mich
Yours GCO.

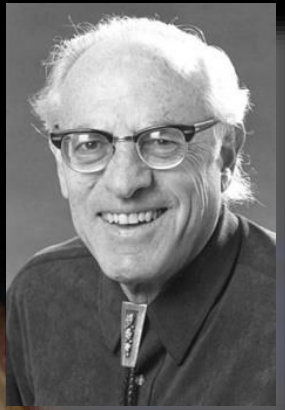


F. Reines answered to G. Gamow:
"Heat loss from Earth's surface is $50 \text{ erg cm}^{-2} \text{ s}^{-1}$.
If assume all due to beta decay than have only enough energy for about 10^8 one-MeV neutrinos cm^{-2} and s."

TO: DR. GEORGE GAMOW
THE UNION
UNIVERSITY OF MICHIGAN
ANN ARBOR, MICHIGAN

MESSAGE:

FROM NUMBERS IN VIKRY BOOK ON THE PLANETS, EQUILIBRIUM HEAT LOSS FROM EARTH'S SURFACE IS $50 \text{ ERGS/CM}^2 \text{ SEC}$. IF ASSUME ALL DUE TO BETA DECAY THEN HAVE ONLY ENOUGH ENERGY FOR ABOUT 10^8 , $1 \frac{1}{2} \text{ Mev}$ NEUTRINOS PER CM^2 AND SEC. THIS IS LOW BY 10^5 OR SO. SHORT HALF LIVES WOULD BE MADE BY COSMIC RAYS OR NEUTRONS IN EARTH. IN VIEW OF RARITY OF COSMIC RAYS: I.E. ABOUT EQUAL TO ENERGY OF STARLIGHT AND OF NEUTRONS IN EARTH THIS SOURCE OF NEUTRONS SEEMS EVEN LESS LIKELY AS A SOURCE OF OUR SIGNAL.



Lesson 1 – Patience is key in geoneutrino research

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nature

ARTICLES

Experimental investigation of geologically produced antineutrinos with KamLAND

After 749.1 ± 0.5 days of detector live-time, the KamLAND experiment measured **25 geoneutrino candidates** from ^{238}U and ^{232}Th decay chains. With geological input adjustments (Th/U mass ratio of 3.9), the estimated event count is $28^{+26.2}_{-23.5}$ at the 90% confidence level.

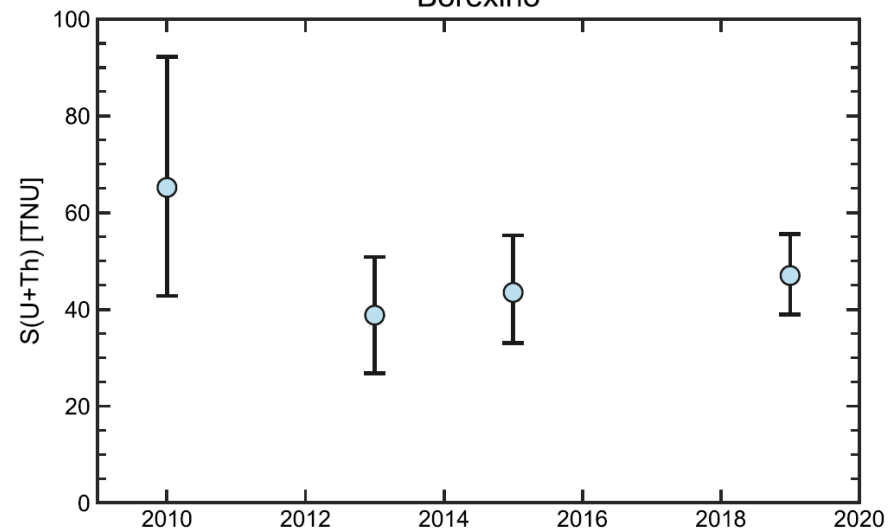


Borexino and KamLAND geoneutrinos results

Borexino

- Period: 2007 – 2019
- Geo- ν events: $52.6^{+7.4}_{-6.3}$
- Signal: $47.0^{+8.7}_{-7.9}$ TNU

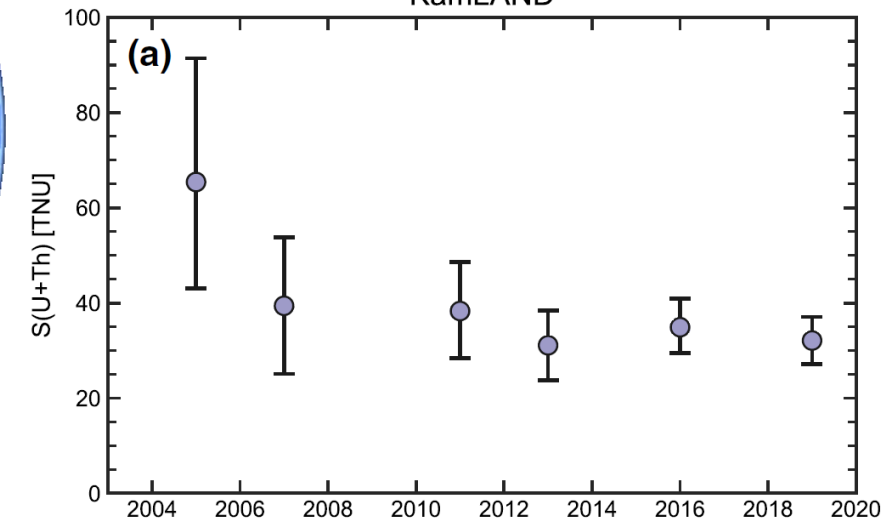
Borexino



KamLAND

- Period: 2002 – 2019
- Geo- ν events: $168.8^{+26.3}_{-26.5}$
- Signal: 32 ± 5 TNU

KamLAND

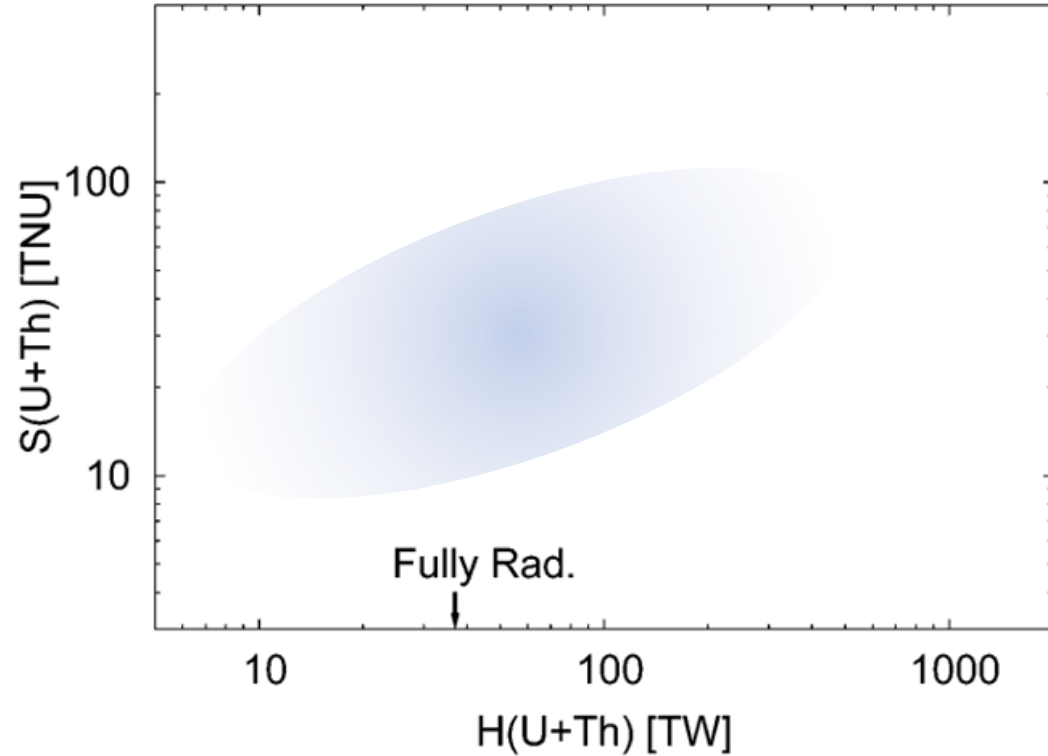
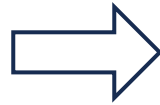
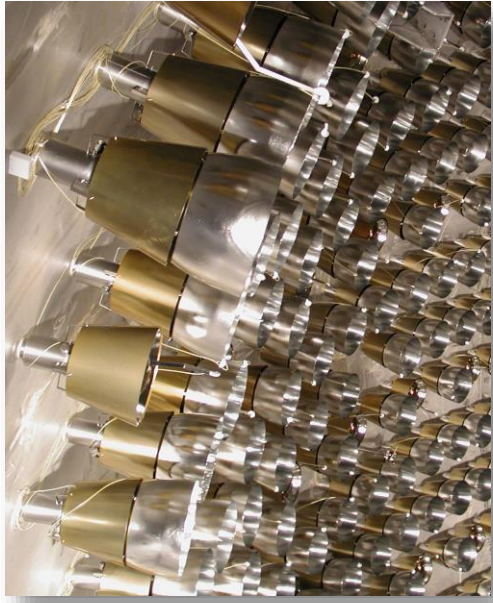


In principle, we have two independent observatories measuring Earth's total radioactivity.

But what can we learn from these measurements?

Lesson 2 – Geoneutrino signals need geological models

Signal is measured in 1 TNU = 1 event per 10^{32} free protons/year



U, Th and ^{40}K in the Earth release heat together with anti- ν , in a well-fixed ratio

Decay	$T_{1/2}$ [10^9 yr]	E_{\max} [MeV]	Q [MeV]	$\varepsilon_{\bar{\nu}}$ [$\text{kg}^{-1}\text{s}^{-1}$]	ε_H [W/kg]
$^{238}\text{U} \rightarrow ^{206}\text{Pb} + 8\ ^4\text{He} + 6e + 6\bar{\nu}$	4.47	3.26	51.7	7.46×10^7	0.95×10^{-4}
$^{232}\text{Th} \rightarrow ^{208}\text{Pb} + 6\ ^4\text{He} + 4e + 4\bar{\nu}$	14.0	2.25	42.7	1.63×10^7	0.27×10^{-4}
$^{40}\text{K} \rightarrow ^{40}\text{Ca} + e + \bar{\nu}$ (89%)	1.28	1.311	1.311	2.37×10^8	0.22×10^{-4}

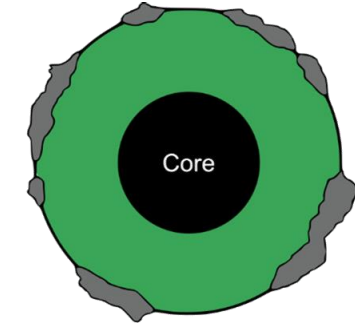


REFERENCE	Earth power
	Q (TW)
Pollack et al., 1993	44 ± 1
Hofmeister and Criss, 2005	31 ± 1
Jaupart et al., 2015	46 ± 2
Davies and Davies, 2010	47 ± 2
Davies, 2013	45
Lucazeau, 2019	44

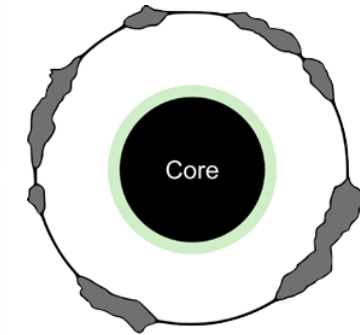
Mantle radiogenic power from KamLAND and Borexino

$$S_M(\text{U+Th}) = \beta H_M(\text{Th+U})$$

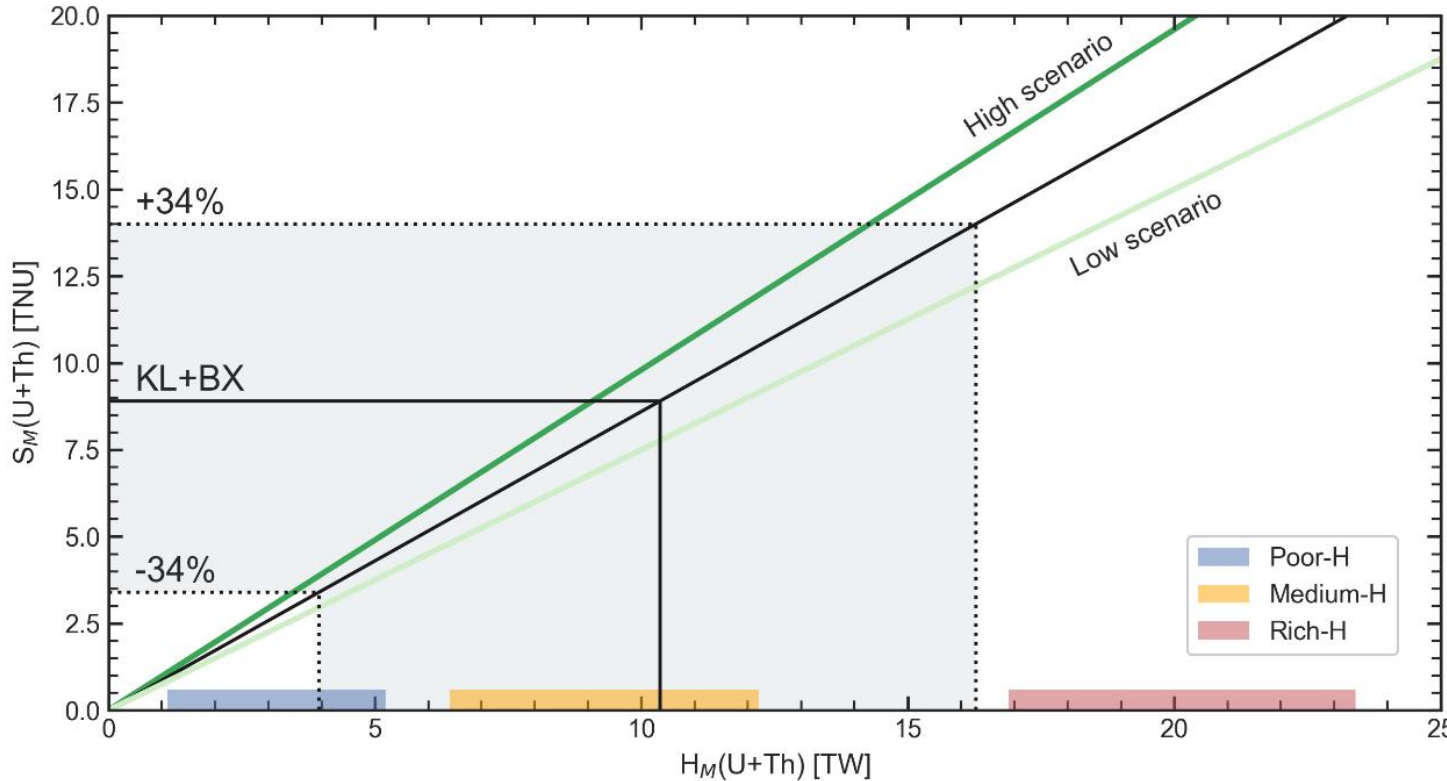
The coefficient β depends only on the distribution of U and Th in the mantle.



$$\beta_{\text{High}} = 0.98 \text{ TNU/TW}$$



$$\beta_{\text{Low}} = 0.75 \text{ TNU/TW}$$



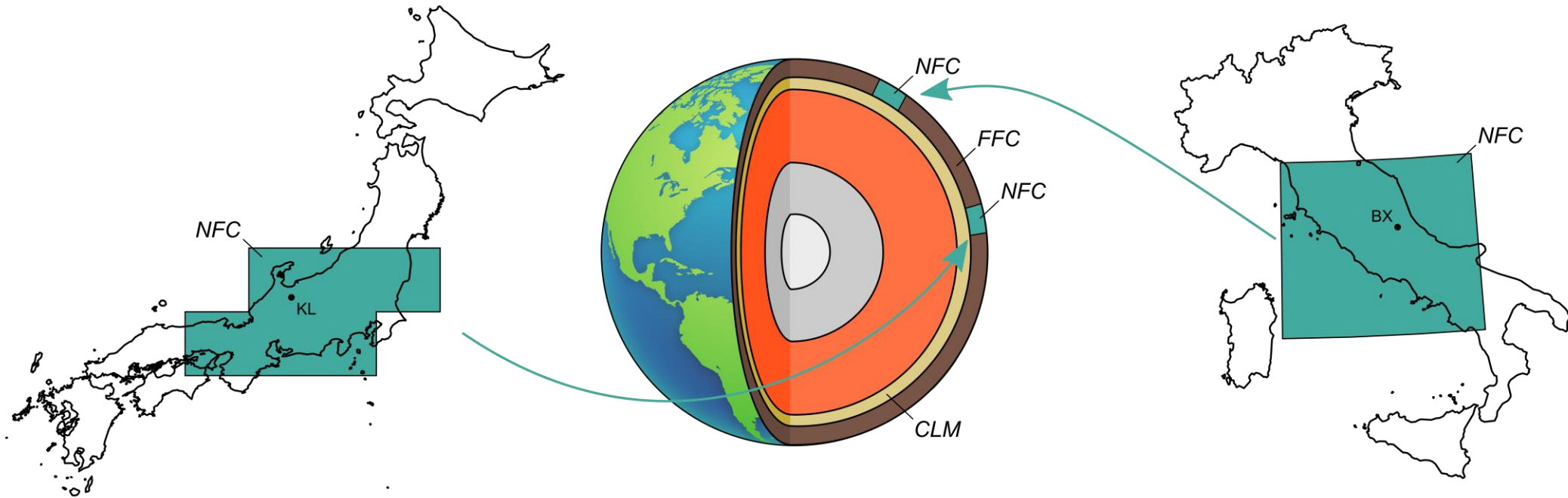
By combining signals from Borexino and KamLAND, a mantle geoneutrino signal of $8.9^{+5.1}_{-5.5}$ TNU* can be extracted

	Geochemical mantle models			
	Poor (U+Th)	Medium (U+Th)	Rich (U+Th)	KL+BX
$H_M(\text{U+Th})$ [TW]	$3.2^{+2.0}_{-2.1}$	9.3 ± 2.9	$20.2^{+3.2}_{-3.3}$	$10.3^{+5.9}_{-6.4}$

* Bellini et al. - Riv. Nuovo Cim. 45, 1–105 (2022)

How can extract the mantle signal from experimental data?

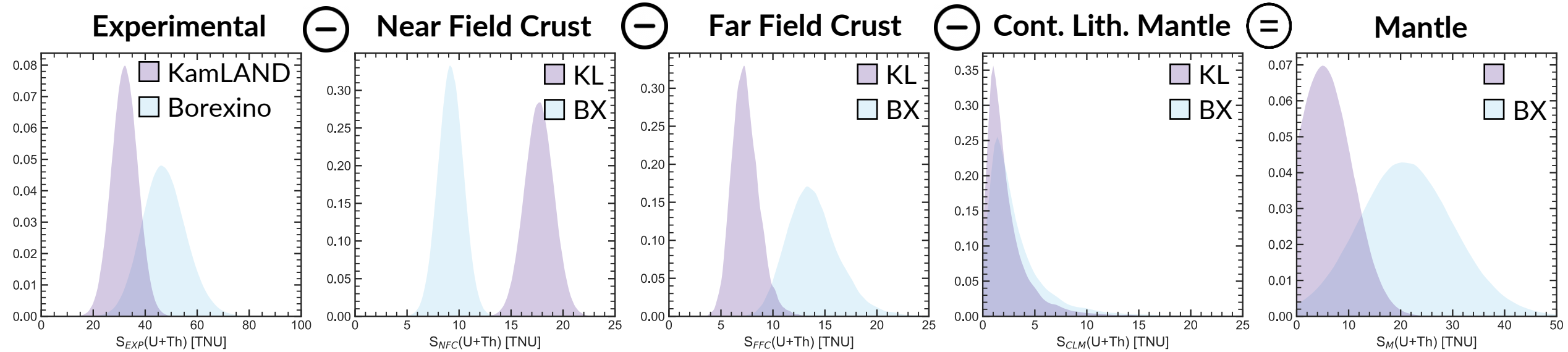
$$S_M(\mathbf{U} + \mathbf{Th}) = S_{\text{Exp}}(\mathbf{U} + \mathbf{Th}) - S_{\text{NFC}}(\mathbf{U} + \mathbf{Th}) - S_{\text{FFC}}(\mathbf{U} + \mathbf{Th}) - S_{\text{CLM}}(\mathbf{U} + \mathbf{Th})$$



- U and Th distributed in the **Near Field Crust (NFC)** (i.e. ~ 500 km within the detector) gives a significant contribution to the signal (~ 50% of the total). The modeling of the **NFC** should be built with local geochemical and/or geophysical information
- The signal of **the Far Field Crust (FFC)** and of the **Continental Lithospheric Mantle (CLM)** is modeling based on global reference models.

Lesson 3 – The better we know the crust, the better we can infer the mantle

The mantle signals $S_M^{BX}(U + Th)$ and $S_M^{KL}(U + Th)$ can be inferred by subtracting the estimated lithospheric components from the experimental total signals using their reconstructed PDFs:

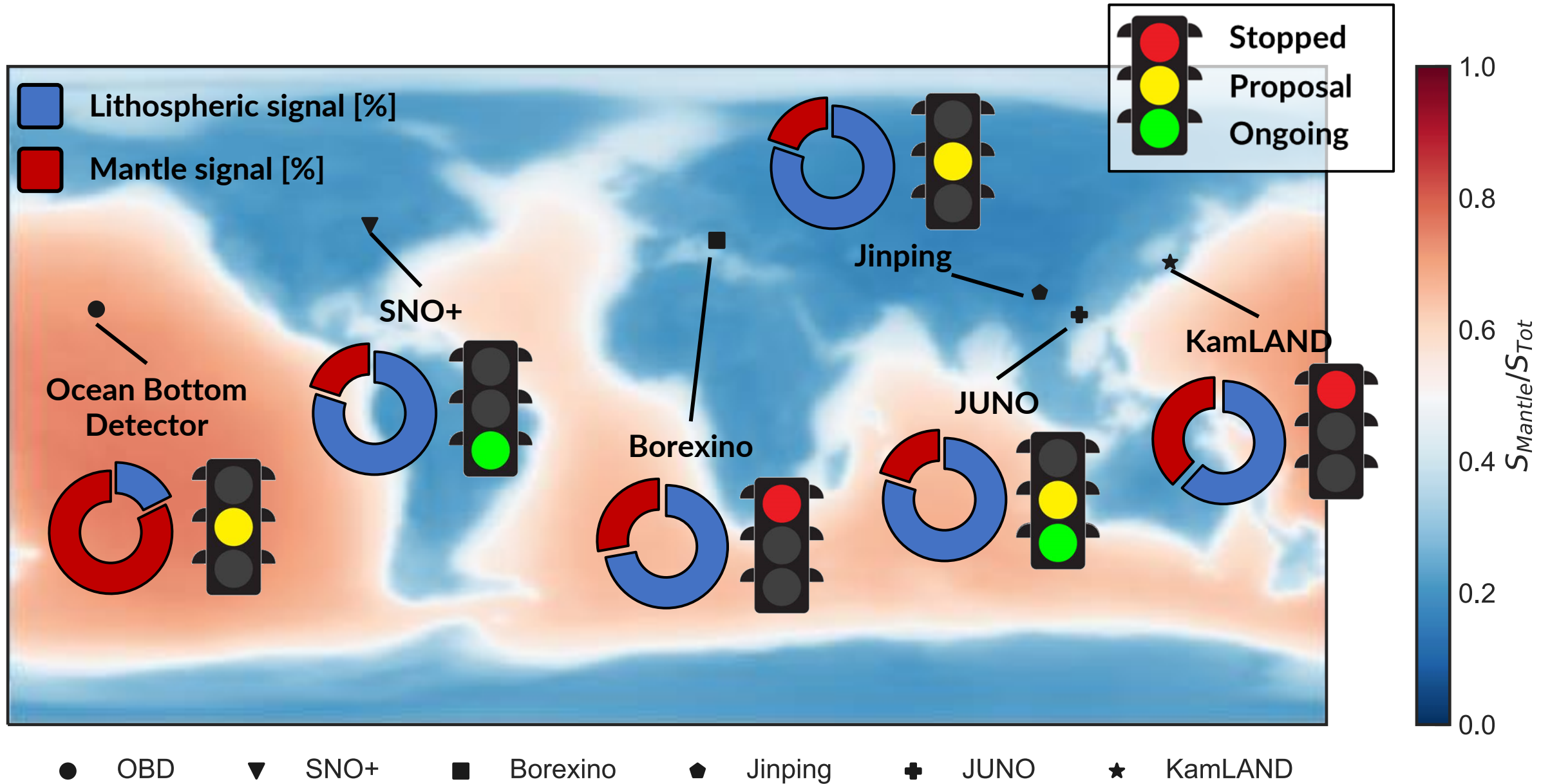


$$S_{Exp}^i(U + Th) - S_{NFC}^i(U + Th) - S_{FFC}^i(U + Th) - S_{CLM}^i(U + Th) = S_M^i(U + Th)$$

$S_{Exp}(U+Th)$ [TNU] $S_{NFC}(U+Th)$ [TNU] $S_{FFC}(U+Th)$ [TNU] $S_{CLM}(U+Th)$ [TNU] $S_M(U+Th)$ [TNU]

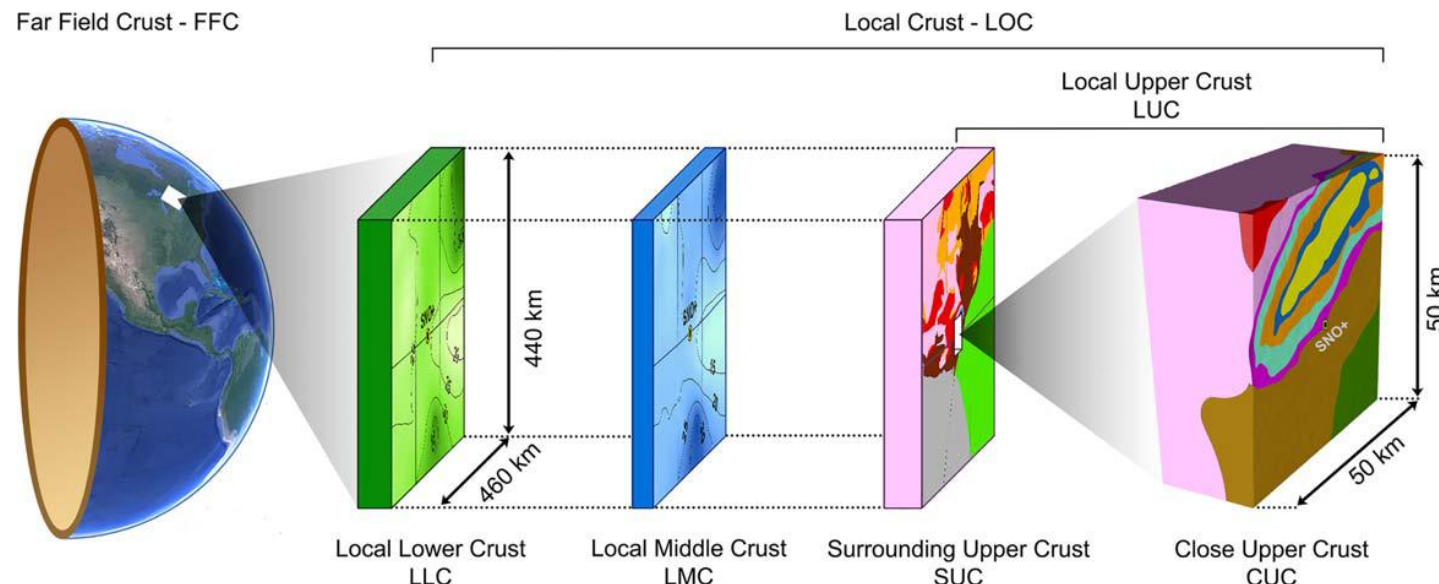
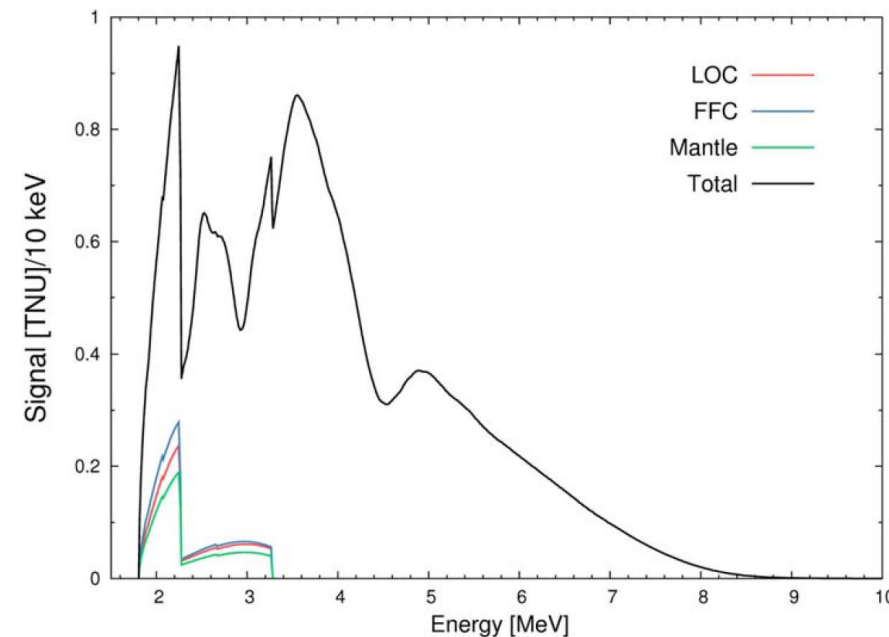
	KL	32.1 ± 5.0	17.7 ± 1.4	$7.3^{+1.5}_{-1.2}$	$1.6^{+2.2}_{-1.0}$	$4.8^{+5.6}_{-5.9}$
	BX	$47.0^{+8.6}_{-8.1}$	9.2 ± 1.2	$13.7^{+2.8}_{-2.3}$	$2.2^{+3.1}_{-1.3}$	$20.8^{+9.4}_{-9.2}$

Present and next generation detectors



Expected geoneutrino signal at SNO+

- In the geoneutrino energy window [1.8–3.3 MeV], an expected reactor antineutrino signal of $48.5^{+1.8}_{-1.5}$ TNU.
- Two distinct 3D models of the SNO+ crust, impacted by a meteorite $\sim 1.9 \cdot 10^9$ yr ago, were developed at different resolutions using geological, geophysical, and geochemical data.
- U and Th in the main 9 upper crust (UC) units were characterized with 112 samples.



	S(U+Th) TNU
LOC	$15.3^{+7.7}_{-3.3}$
FFC	$15.2^{+2.7}_{-2.4}$
CLM	$2.1^{+3.0}_{-1.3}$
Lithosphere	$34.2^{+9.2}_{-5.3}$

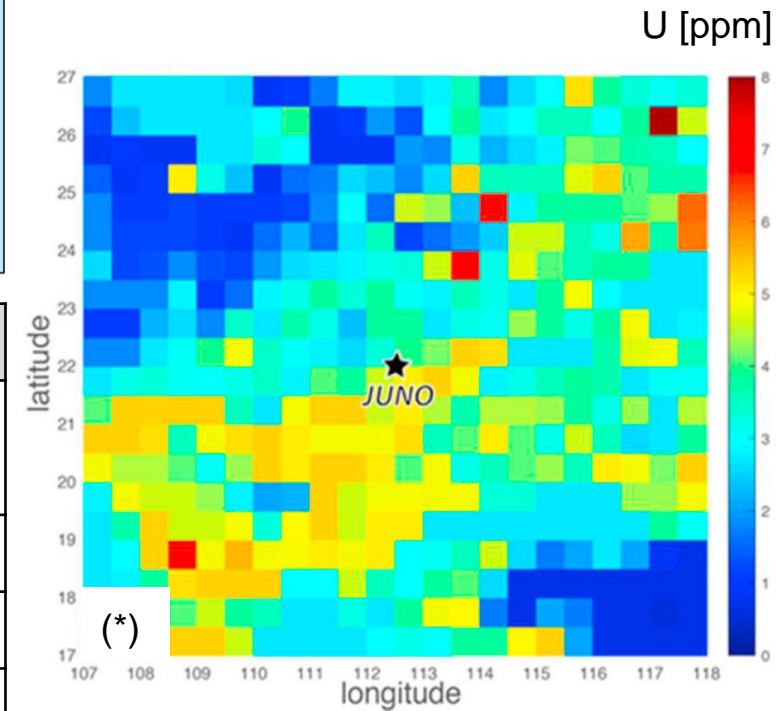
The total expected geoneutrino signal at SNO+ is estimated at 42 ± 8 TNU

Expected geoneutrino signal at JUNO

- JUNO is a 20 kton LS detector surrounded by ~18.000 20" PMT
- Expected geo- ν ~ 400 events/year (~ 40 TNU)
- Expected react- ν in [1.8-3.3 MeV] ~ 260 TNU ($S_{\text{rea}} / S_{\text{geo}} \sim 7$)



	N° of cores	Thermal power/core
Yangjiang	6	2.9 GW
Taishan	2	4.6 GW



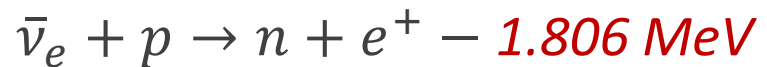
	S(U+Th) [TNU]
Strati et al., 2015 (using global crustal model)	$39.7^{+6.5}_{-5.2}$
Wipperfurth et al., 2020 (using global crustal models)	$41.3^{+7.5}_{-6.3}$
	$41.2^{+7.6}_{-6.4}$
	$40.0^{+7.4}_{-6.2}$
Gao et al., 2020 (*) (combining global crustal model and local geological data)	$49.1^{+5.6}_{-5.0}$

...1953 -> 2024



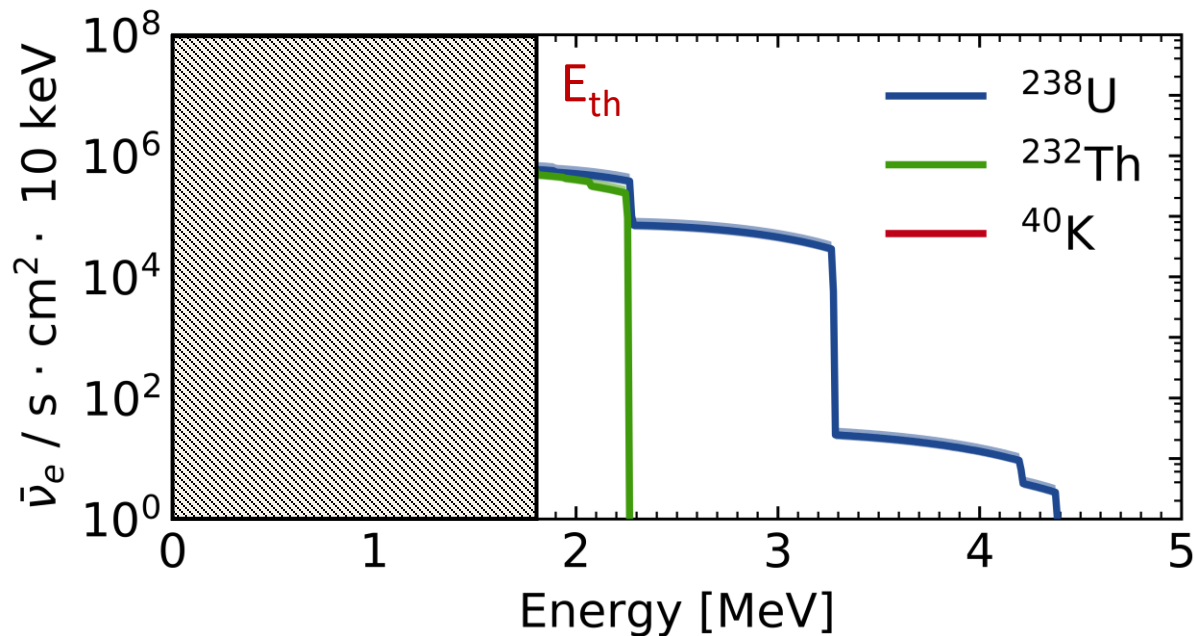
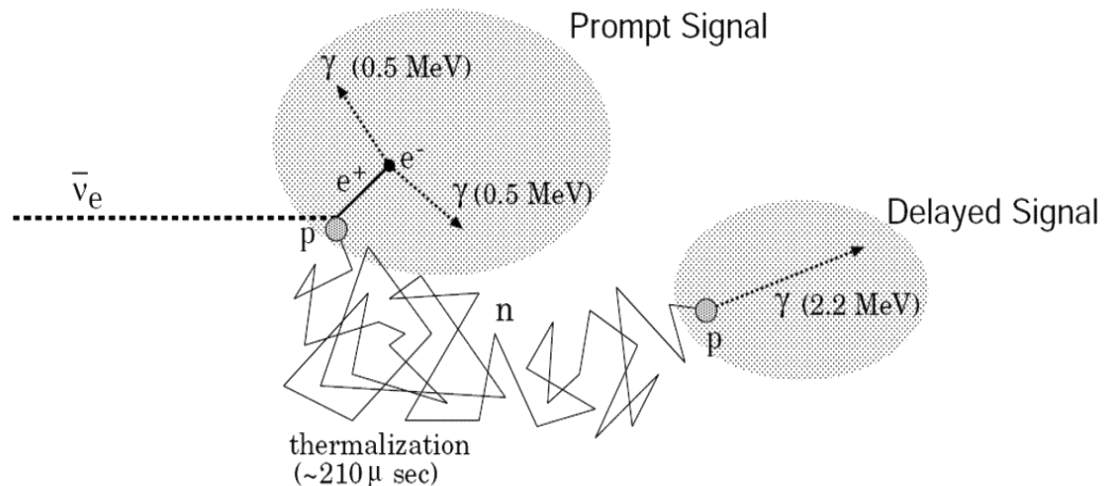
Inverse Beta Decay (IBD) detection

Geoneutrinos are **detected by IBD** in \sim kton Liquid Scintillation Detectors.



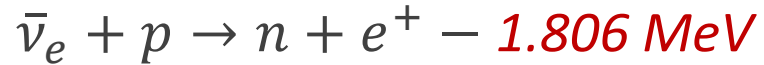
Detection requires the coincidence of 2 delayed light signals.

It does not permit to observe $^{40}\text{K}-\bar{\nu}_e$



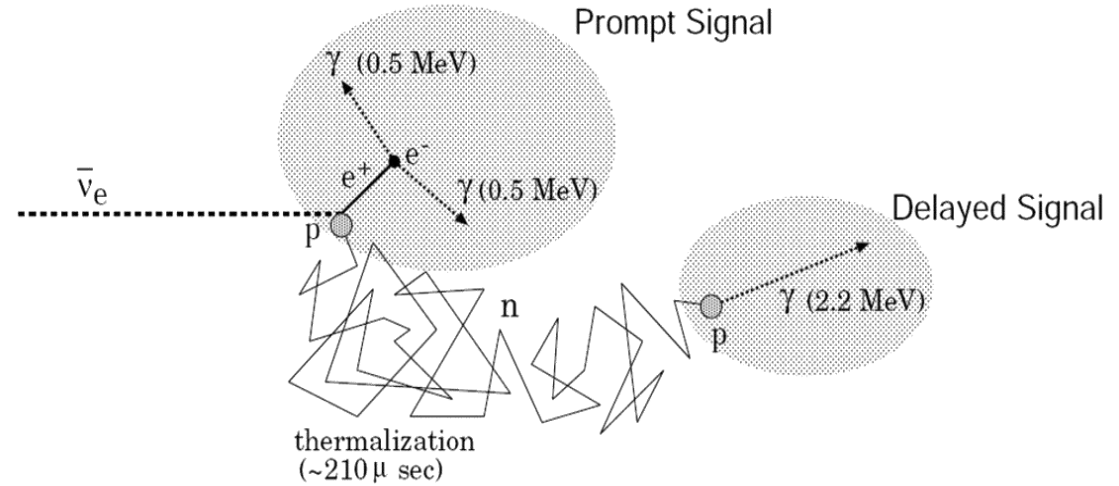
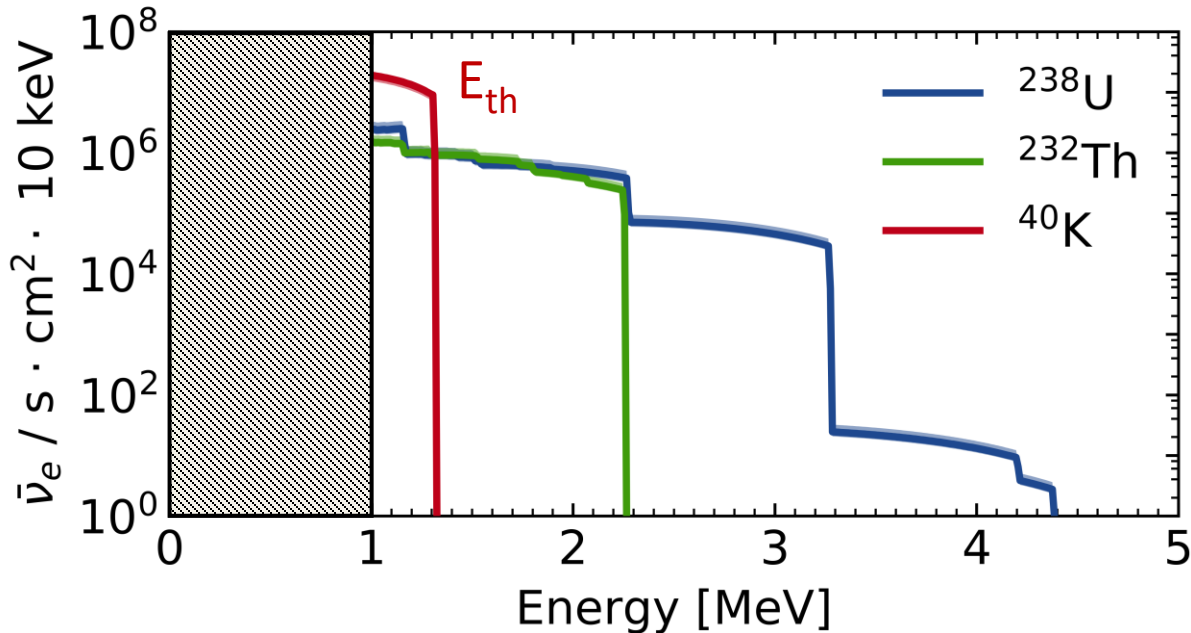
Inverse Beta Decay (IBD) detection

Geoneutrinos are **detected by IBD** in \sim kton Liquid Scintillation Detectors.

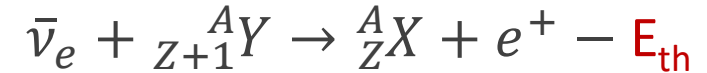


Detection requires the coincidence of 2 delayed light signals.

It does not permit to observe $^{40}\text{K}-\bar{\nu}_e$



In order to detect $^{40}\text{K}-\bar{\nu}_e$ we could use:

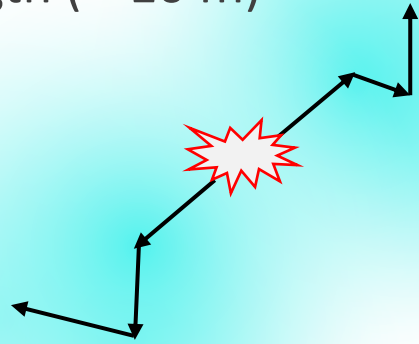


We shall require:

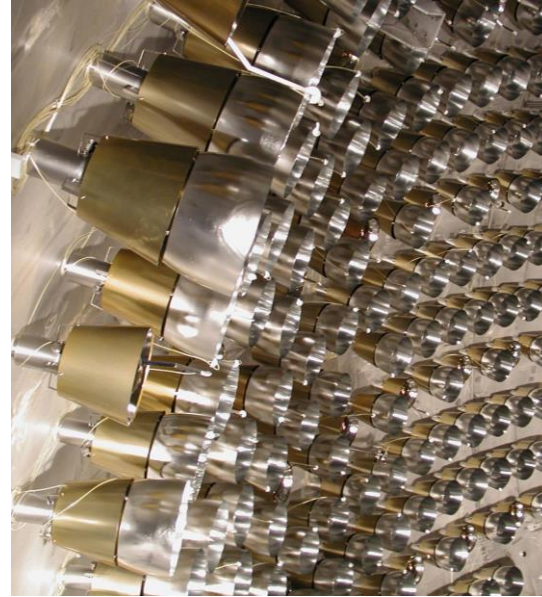
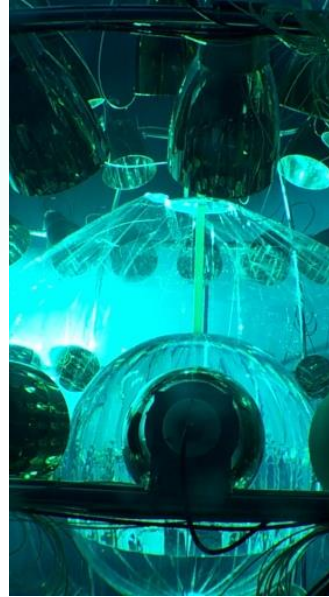
- $E_{th} < 1.3 \text{ MeV}$
- High cross-section
- High Y natural isotopic abundance

Transparent vs. opaque detector

Very long scattering length (~ 10 m)

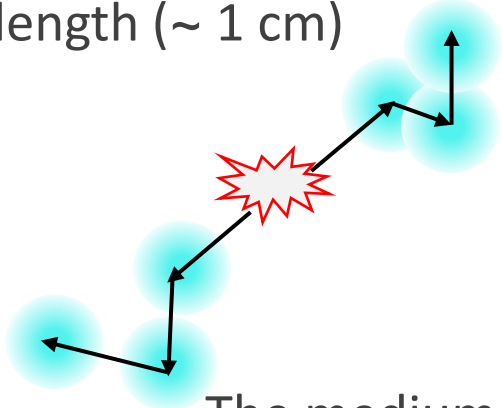


The medium is transparent to scintillation photons

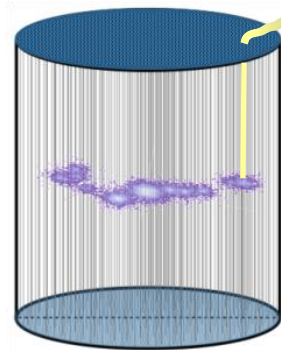
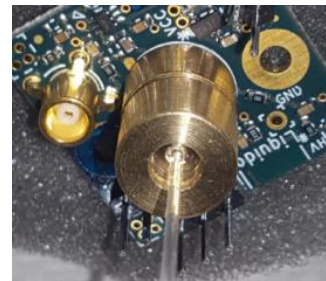


- Scintillation light reaches the surrounding 10^3 - 10^4 PMTs
- Slow time resolution (\sim ns)
- Poor spatial resolution on light deposition (~ 10 cm)
- High photon detection efficiency ($\sim 20\%$)

Very short scattering length (~ 1 cm)

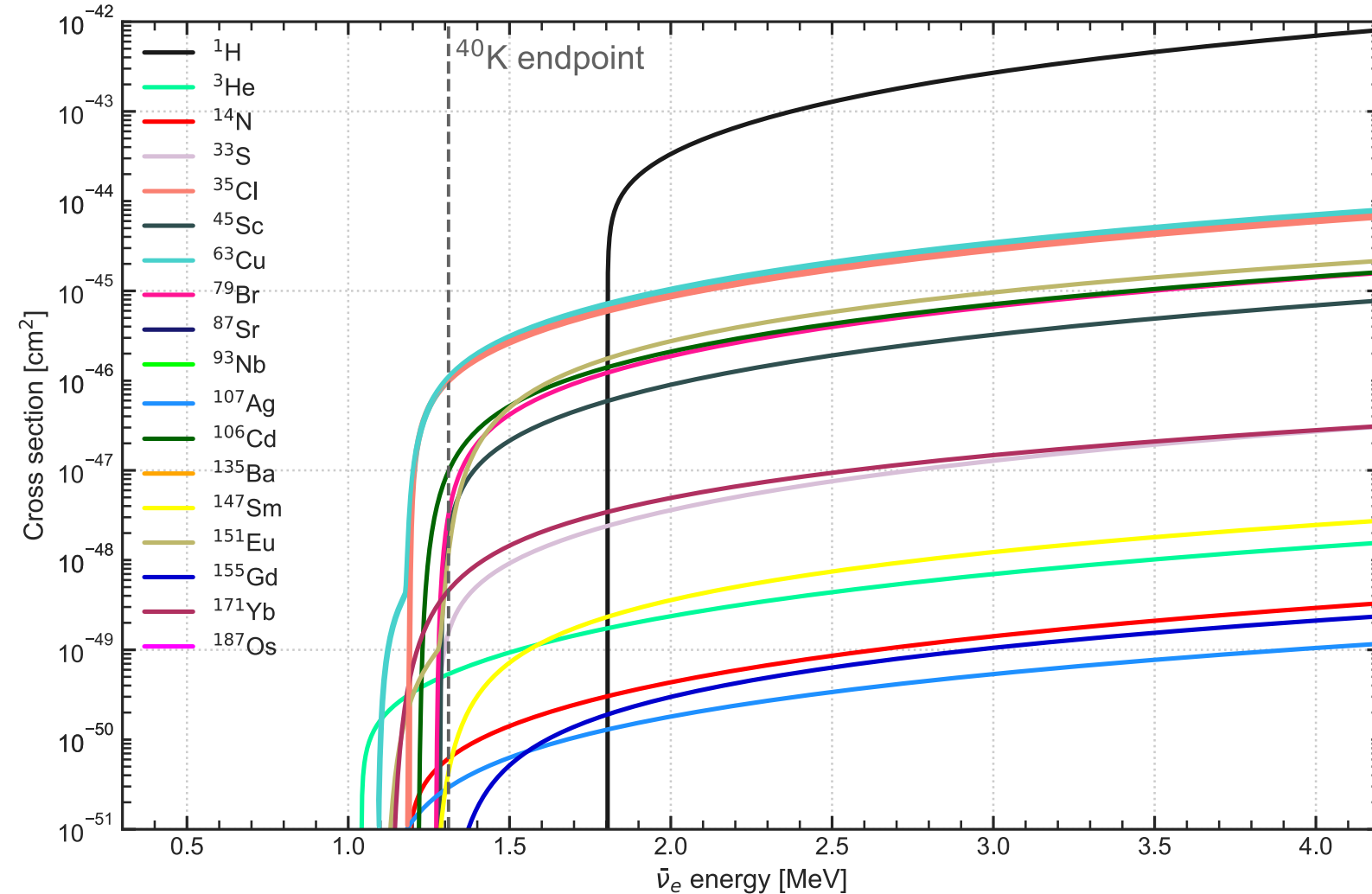


The medium is opaque to scintillation photons

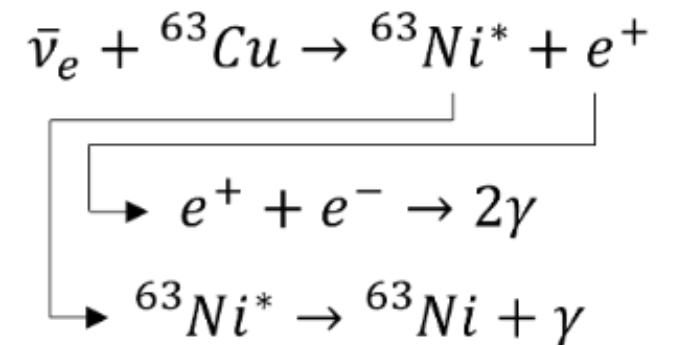


- The light is extracted by an array of optical fibers connected to SiPMs
- Fast time resolution (~ 0.3 ns)
- Excellent spatial resolution on light deposition (~ 1 cm)
- Poor photon detection efficiency ($\sim 5\%$)

IBD cross-sections weighted by isotopic abundance



^{63}Cu (Isotopic Abundance = 69%) appears to be a promising target for ^{40}K geoneutrinos due to its transition to an excited state in $^{63}\text{Ni}^*$ ($E_{\text{MAX}} = 1.176$ MeV; $t_{1/2} = 1.67$ μs), offering potential double-coincidence capability ($E_\gamma = 87$ keV).





Conclusions & perspectives

- We learn three lessons from past research:
 - patience is essential in geoneutrino studies;
 - interpreting geoneutrino signals requires geological models;
 - the better we understand the crust, the more accurately we can infer mantle properties.
- The "multi-site detection" era for U+Th geoneutrinos is underway, with new data from SNO+ and JUNO soon reinforce constraints on mantle composition and Earth's radiogenic heat.
- The development of new technologies (i.e. opaque liquid-based detectors) may open opportunities for detecting potassium geoneutrinos.

